Project presentation

- Topic has to be chosen before 31 May
- Start early → topics need refinement
- Presentation of 15 mins total
 - 12 mins talk + 3 mins Q&A
- Presentations in English
- Evaluation criteria:
 - Clarity, creativity, integrity
 - Results not important
- > Two people pair up for a project
 - Joint report with author contribution statement
 - Joint presentation

WORKLOAD

- > 3 ECTS Course
- > 90 h in total for the entire semester
 - 90 h = 30 h lecture + 60 h self study
- > 6 h per week for 15 weeks
 - > 6 h = 2 h lecture + 4 h self study
- Division of self study workload (60h)
 - Per person:
 - Lecture comprehension: 15 h
 - Project work: 30 h
 - Project presentation & report: 15 h
 - Per project team (of two members):
 - Project work: 60 h
 - Project presentation & report: 30 h

LECTURE 2: Introduction to the earth's climate

ML-4430: Machine learning approaches in climate science

28 Apr 2021



- Weather, on average?
- > The climate system

Γhe Atmosphere



- Vertical structure
- Meridional Structure

Components of the climate system

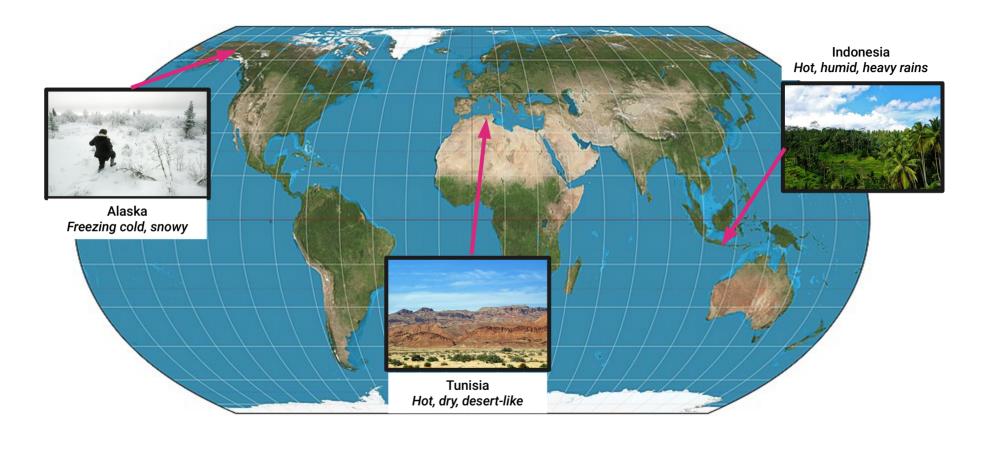


- Atmosphere
- > Ocean
- > Land
- Cryosphere
- Biosphere
- Anthroposphere



- Climate Forcings
- Climate Feedbacks
- Climate Sensitivity

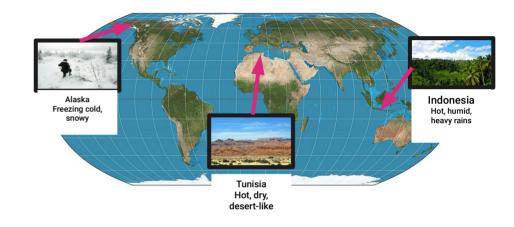




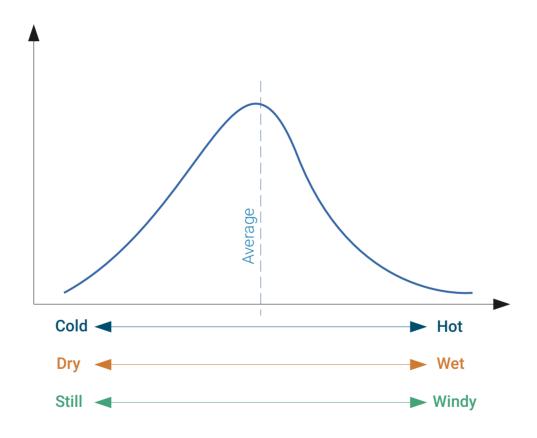


Several questions arise:

- Is the weather the same all year around?
- Does it differ depending on time of day?
- Do certain weather events (droughts, floods, blizzards, etc.) repeat at longer intervals such as decades or centuries?
- What about longer time scales?



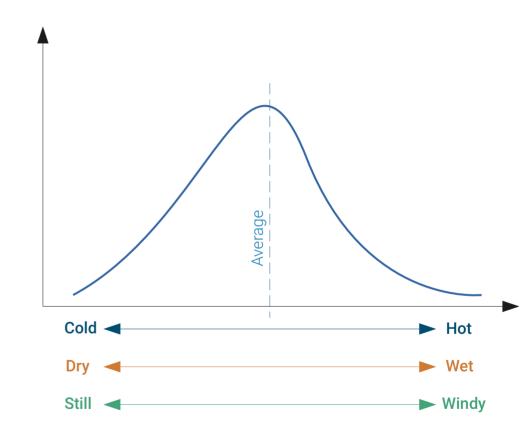




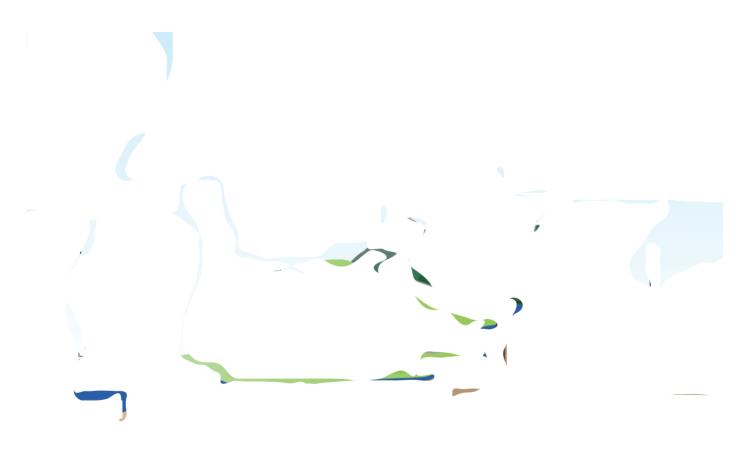


Climate is a distribution of (weather) states

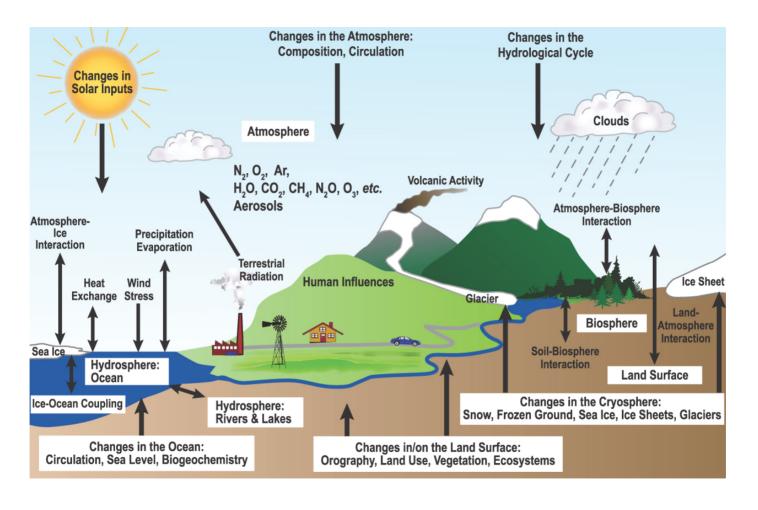
- Given a particular location or region
- Given a particular temporal scale (daily, weekly, annual, decadal, etc.)
- Climate change is thus a change in the distribution of weather states
- A change in the mean is only one (of many) indicator of climate change















The Atmosphere

Vertical structureMeridional Structure

Components of the climate system



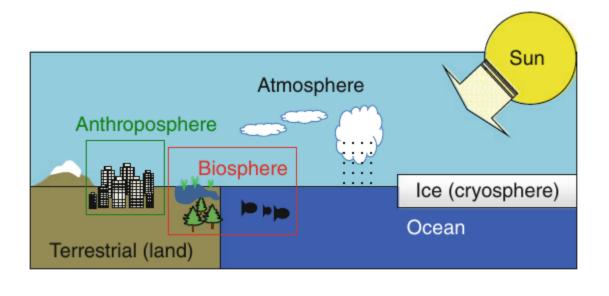
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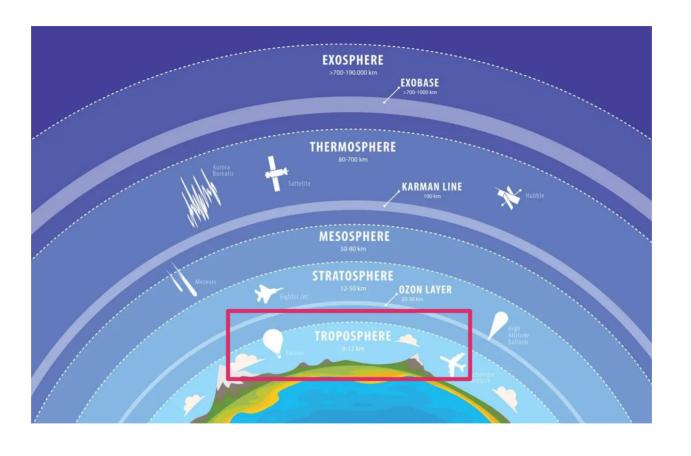
- Climate Forcings
- Climate Feedbacks
- Climate Sensitivity



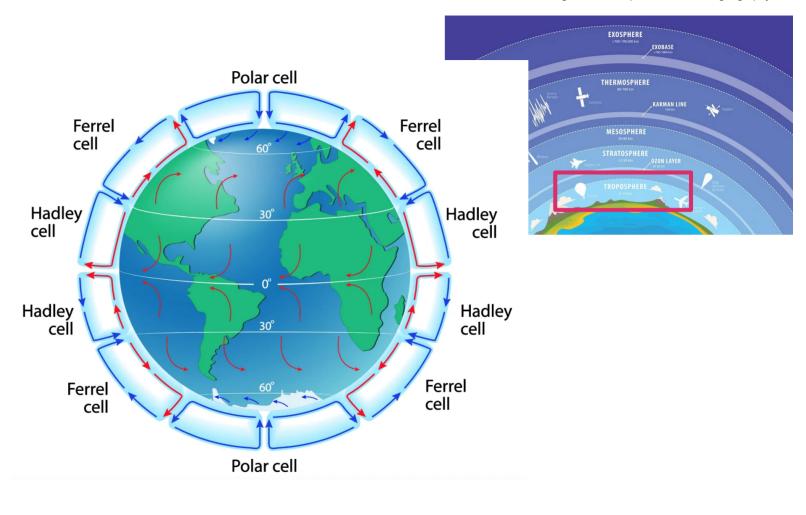








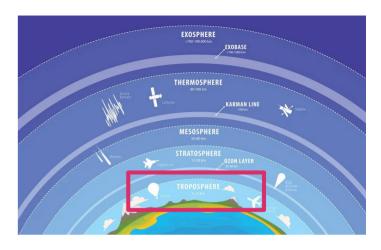


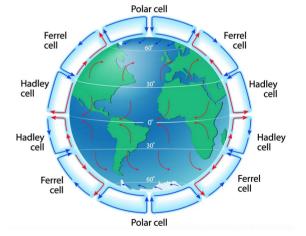




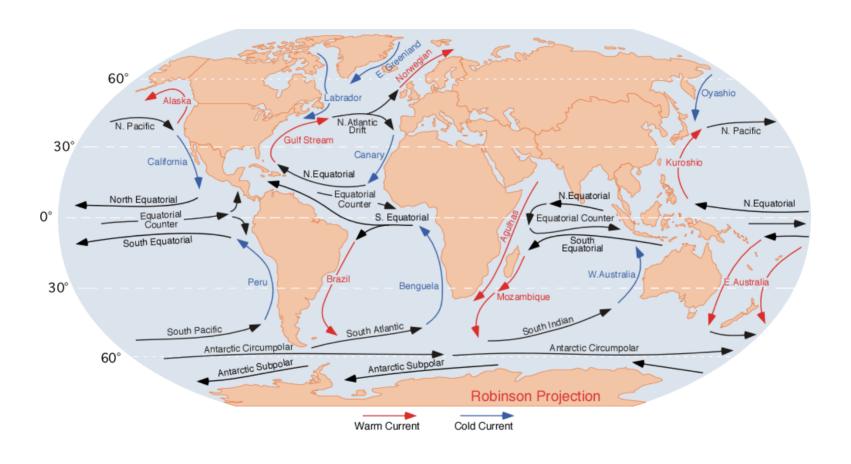
Atmosphere

- Most relevant climate dynamics in the troposphere
- Exchange of heat and energy with land, ocean, cryosphere, and the sun
- Marked by dominant and stable patterns of wind circulation
- Greenhouse gases: Water and CO2 → both are cycled through the atmosphere

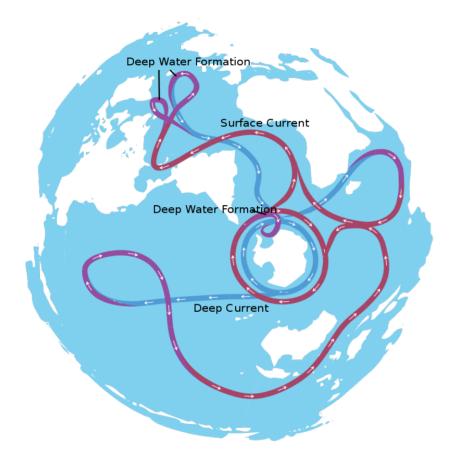


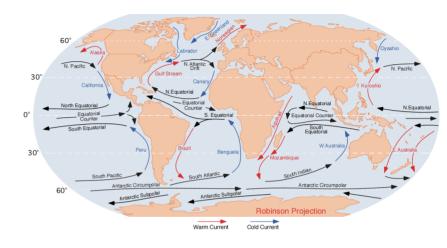








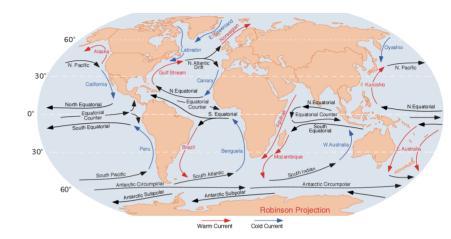


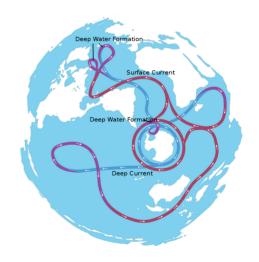




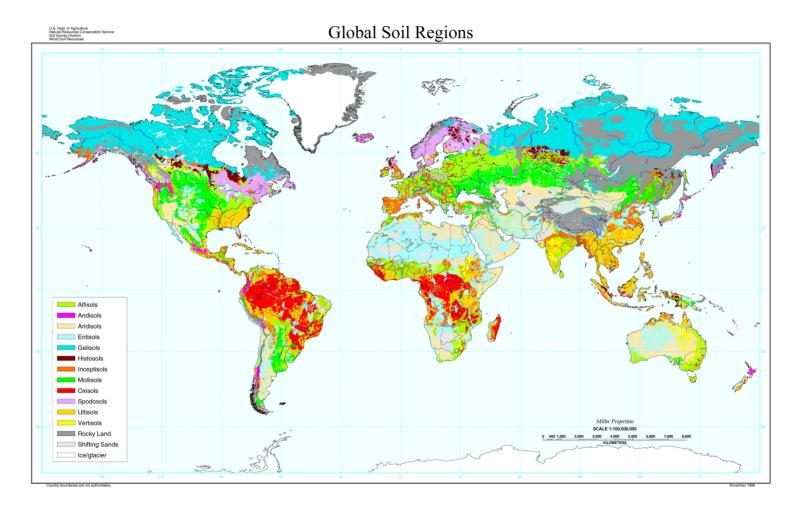
Ocean

- Dynamics composed of surface currents and deep water (thermohaline) currents
- Surface currents forced by earth's rotation and surface winds
- Thermohaline circulation forced by density (and thus salinity) and temperature differences
- Dominant mode of northward transport of heat in the Atlantic

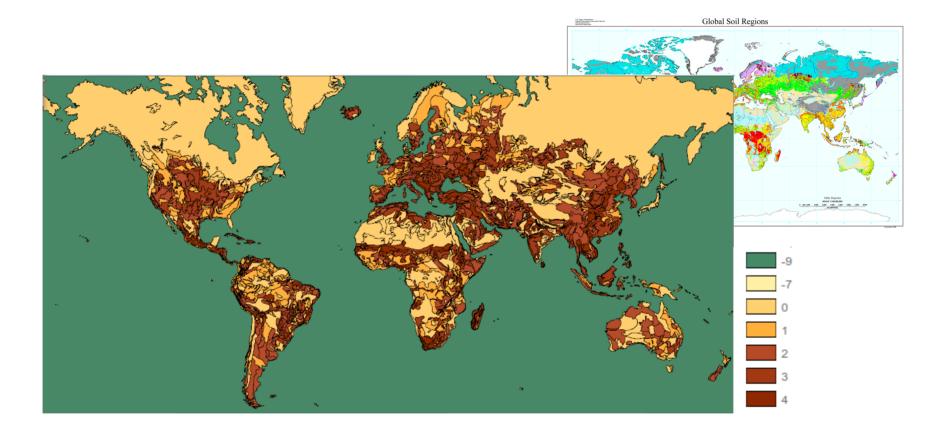








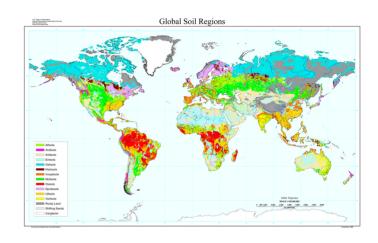


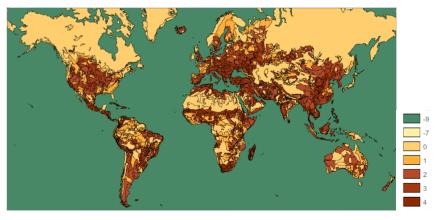




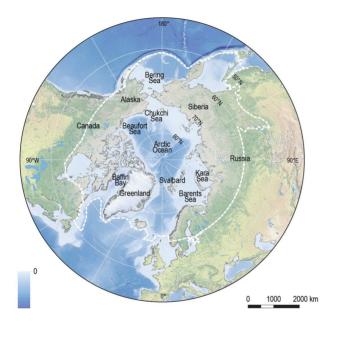
Land

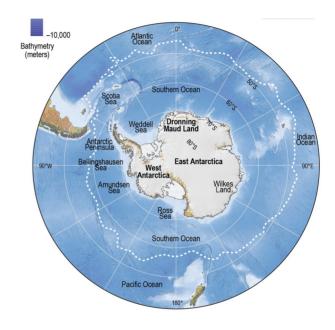
- Soil plays the most important role
- Modulates water, CO2, and other important cycles
- Crucial determinant in climatological conditions
- Interface of atmosphere, cryosphere, ocean, and biosphere



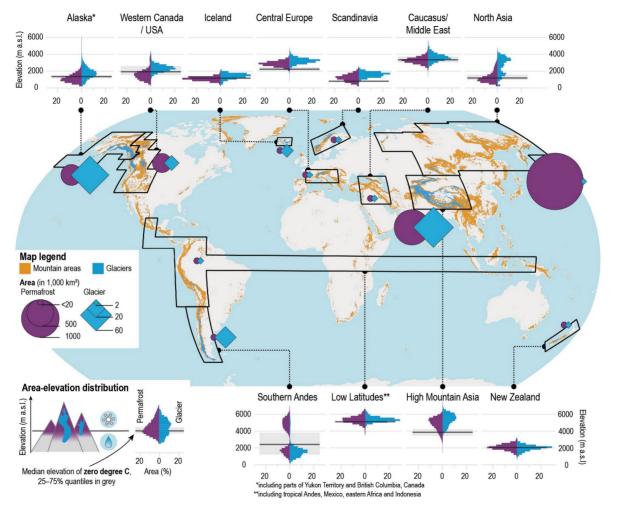












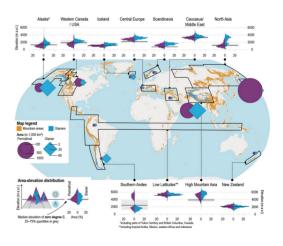


Cryosphere

- Polar regions and high mountain areas
- Important source of freshwater
- Global ice extent modulates global climate
- Arctic and Antarctic ice extent crucial to maintaining polar fronts which influence extratropical climate
- Glaciers influence regional and hemispherical climate systems











- Weather, on average?
- > The climate system

The Atmosphere



- Vertical structure
- Meridional Structure

Components of the climate system



- > Atmosphere
- > Ocean
- > Land
- Cryosphere
- Biosphere
- Anthroposphere



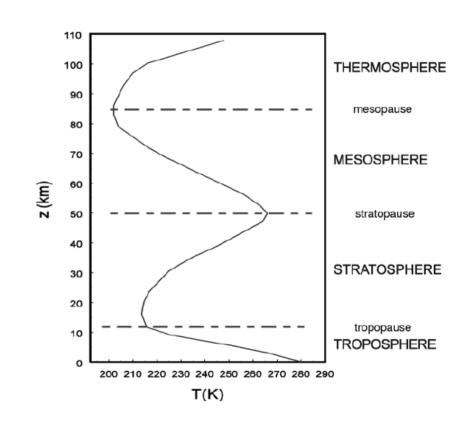
- Climate Forcings
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- Climate Sensitivity





Vertical temperature profile

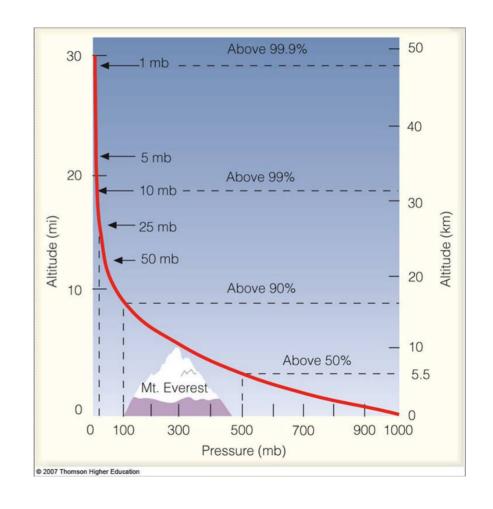
- 3 regions of warm temperature: surface, stratopause, and thermosphere
- Short wavelength UV absorbed in thermosphere (can get really hot)
- Medium wavelength UV absorbed in stratopause (by ozone)
- Stratosphere is stratified; little to no mixing
- Troposphere has all the water vapour; 85% mass of the entire atmosphere





Vertical pressure profile

- Decreases monotonically with altitude
- Pressure at surface is at 1000 mb (=hPa) by convention
- Tropopause (approx 10-16 km) is at around 100 – 200 mb pressure
- Allows a stratification of the atmosphere according to pressure levels (850 mb height, 250 mb height, and so on)
- Possible due to hydrostatic balance





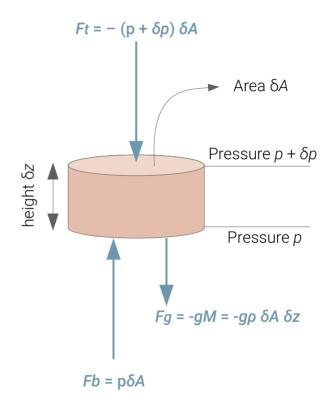
Hydrostatic balance

 Assume that there is no vertical acceleration, i.e. net force on the cylinder of air is zero

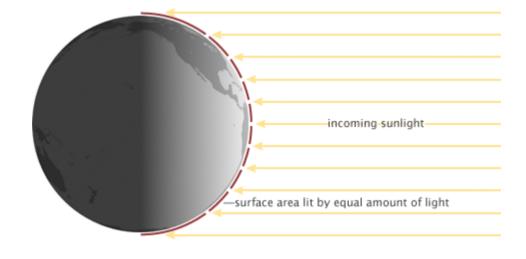
Ft + Fb + Fg = 0

$$\Rightarrow$$
 - $(p + \delta p) \delta A + p \delta A - g \rho \delta A \delta z = 0$
 \Rightarrow $\delta p / \delta z = -g \rho$

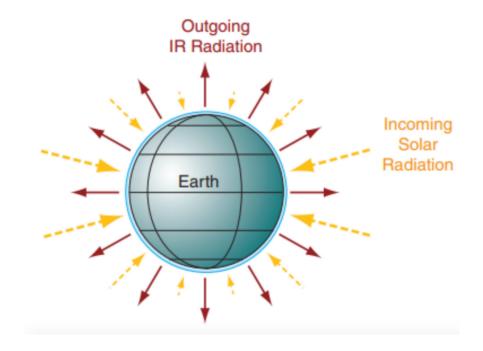
- Pressure decreases with height in proportion to the weight of the overlying atmosphere
- The atmosphere fades away ...

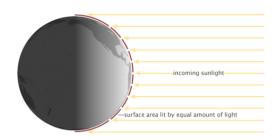




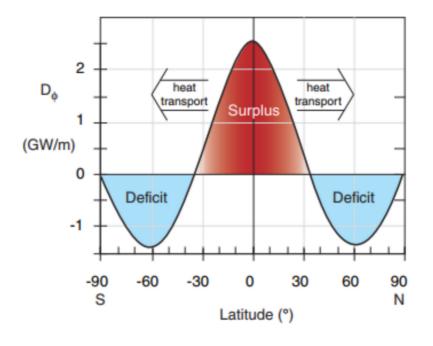


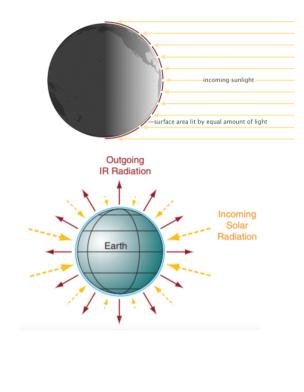




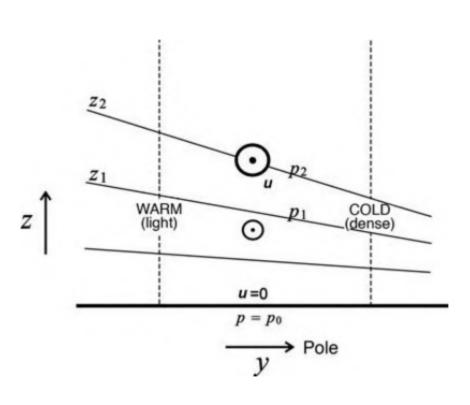


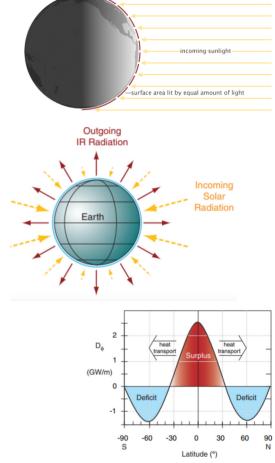








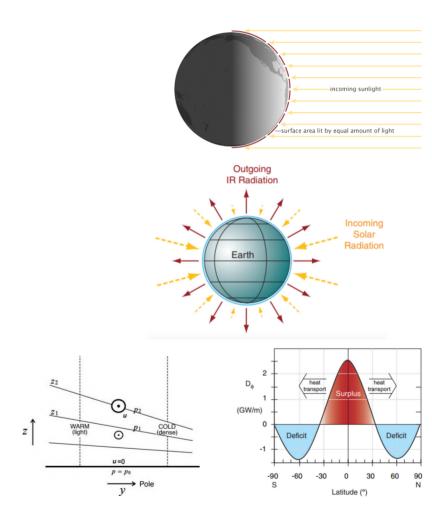






Meridional temperature and pressure gradient

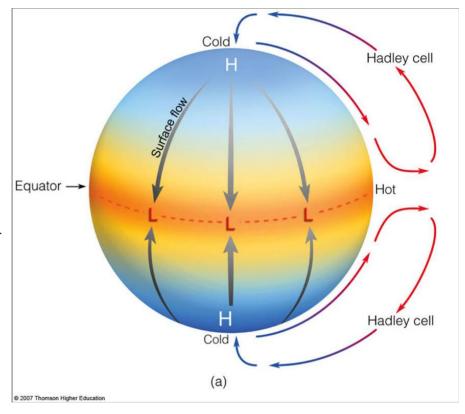
- Sun heats up the equators more than the poles (due to the angle of inclination of sun rays)
- Outgoing IR radiation is more or less even
- Net radiation budget is such that there is a deficit at the poles and a surplus at equator
- Equator air is thus warmer, lighter and taller
- Polar air is thus colder, denser, and shorter
- Meridional pressure gradient is established





Single Cell Model

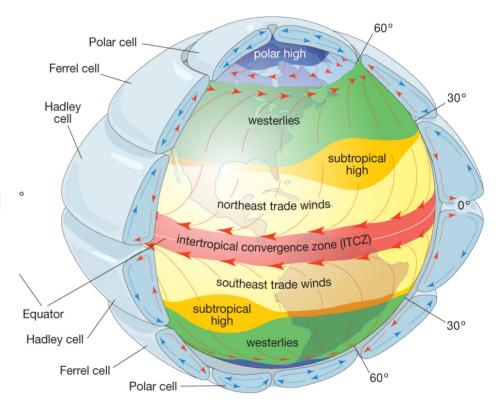
- Assumptions:
 - No land masses
 - No seasons
 - No rotation
- Thermally direct circulation: Hadley Cell
 - "thermally direct": warm air rises and cold air sinks
- In presence of rotation, air is deflected eastwards due to Corioli's Force
- Air masses from equator never make it to the poles!





Three Cell Model

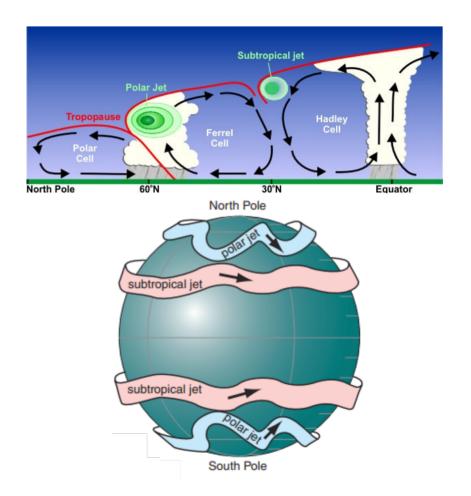
- Poleward moving air from the equator sinks at around 30N and 30S
 - This "subsidence" of dry air coincides with most deserts
- The subsiding air is again sent equatorward and poleward
 - The northeasterly and southeasterly "trade winds" converge around the equator → intertropical convergence zone
 - Poleward westerlies converge at the polar front at 60N and 60S





Jet Streams

- Westerly jets in the upper troposphere (around 250 hPa 150 hPa)
 - Subtropical jet: 30N and 30 S
 - Polar jet: 60N and 60S
- Due to energy imbalance on either side of polar and subtropical fronts respectively
- Strong pressure gradient across the fronts causes intense wind speeds
- The subtropical jets do not extend till the surface







The Atmosphere

Vertical structure

Meridional Structure

Components of the climate system



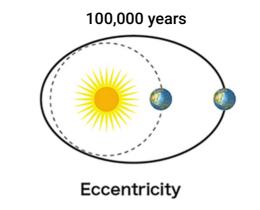
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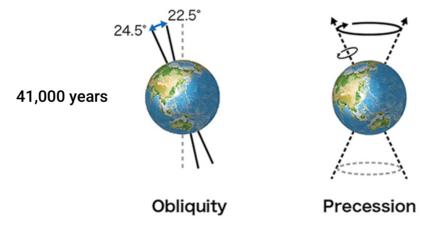


- Climate Forcings
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- Climate Sensitivity



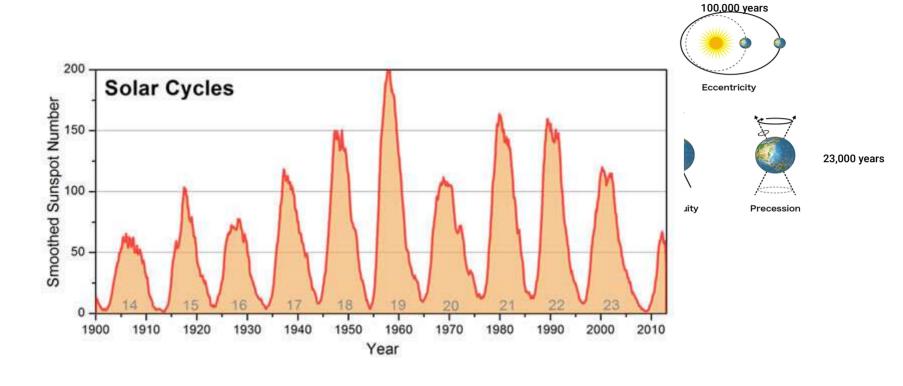






23,000 years

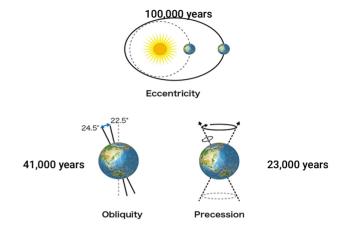


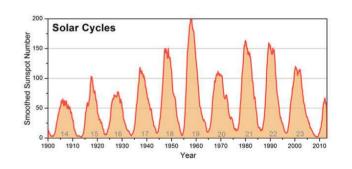




Climate forcings: External

- Orbital forcings, aka Milankovitch forcings
 - Changes in the earth's orbit changes the incoming solar radiation budget
 - Leads to changes in the climate
- Solar cycles, i.e., periodic ups and downs in the sun's activity leads to changes in incoming solar radiation, leading to changes in the climate







Climate Forcings: Internal

- Anthropogenic
 - Greenhouse gases
 - Stratospheric ozone
 - > Tropospheric aerosols
 - Land surface change
- Natural
 - Volcanic aerosols
 - Thermohaline circulation



Climate sensitivity

- Change in a climate parameter for a given change in a climate forcing
 - > Typical example: What is the climate sensitivity of temperature with respect to a doubling of atmospheric CO2?
- Two kinds of climate sensitivity:
 - Transient Climate Response
 - Equilibrium Climate Sensitivity
- Estimated from:
 - Historical climate records
 - Climate models
 - Paleoclimate records





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Q&A